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Modelling of Soil Organic Carbon dynamics in wetlands

Joint work of V. Bohaienko, F. Diele, C. Marangi, A. Martiradonna, A. Provenzale



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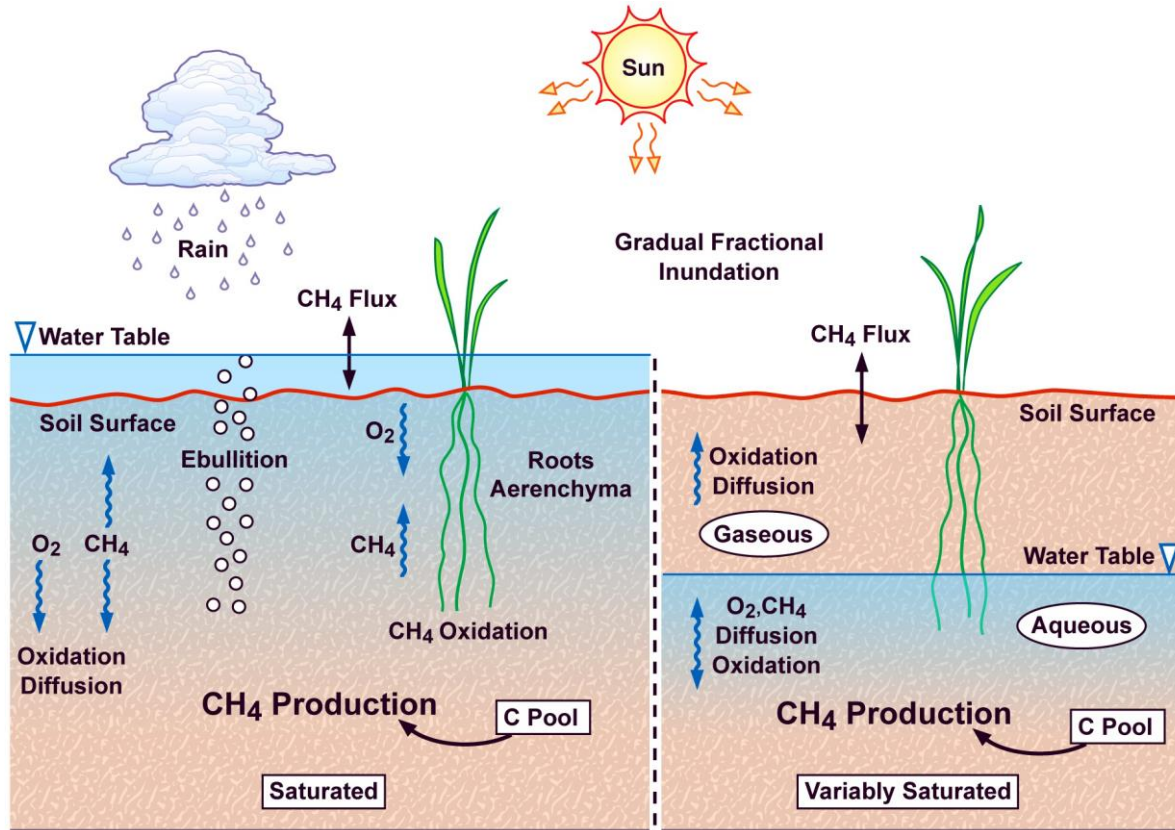
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Modelling of Soil Organic Carbon Dynamics in Wetlands



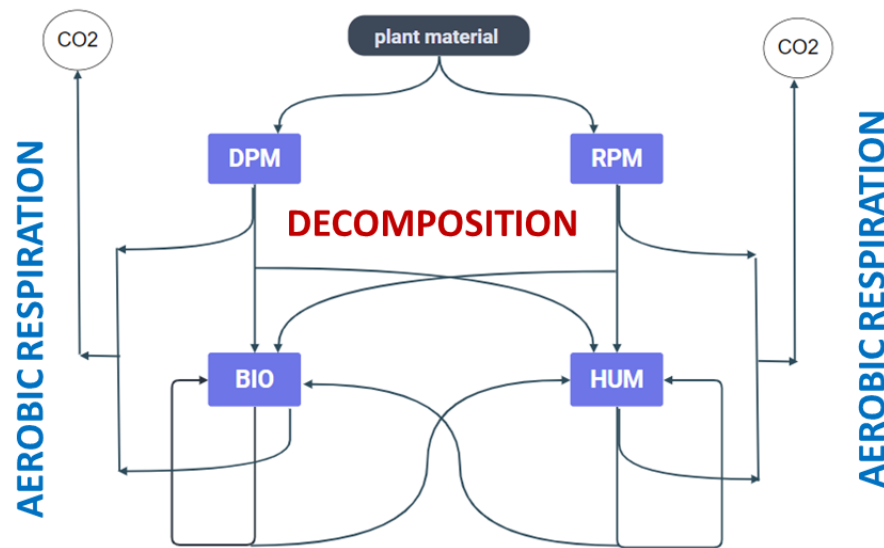
Riley, W. J., Subin, Z. M., Lawrence, D. M., Swenson, S. C., Torn, M. S., Meng, L., ... & Hess, P. (2011). Barriers to predicting changes in global terrestrial methane fluxes: analyses using CLM4Me, a methane biogeochemistry model integrated in CESM. *Biogeosciences*, 8(7), 1925-1953.

RothC Model



Model's assumption:

In unsaturated zones of soil SOC decomposition is performed by bacteria that respire CO₂

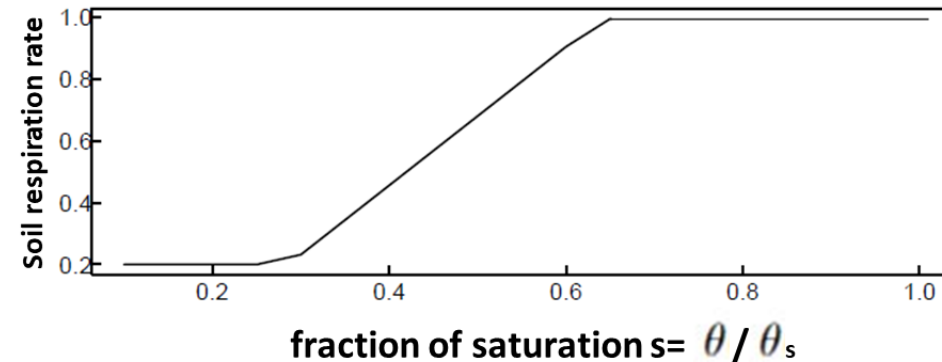


Diele, F., Marangi, C., & Martiradonna, A. (2021). Non-standard discrete RothC models for soil carbon dynamics. *Axioms*, 10(2), 56.

Carbon decomposition is modeled by:

Intrinsic respiration rates that describe the transition from bacteria activity resulting in CO₂ generation

Modifying rate factor for soil moisture

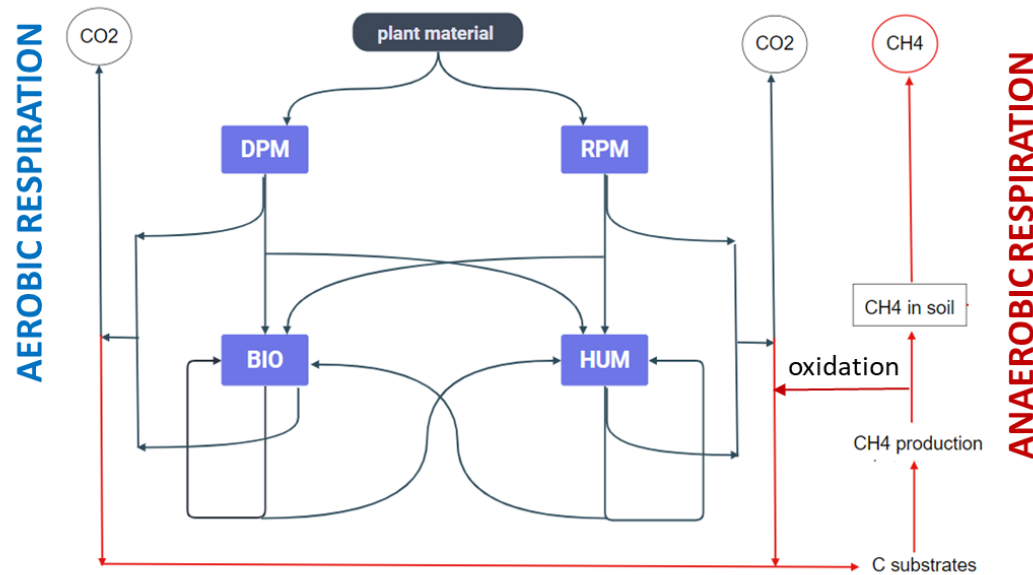


Model's assumptions

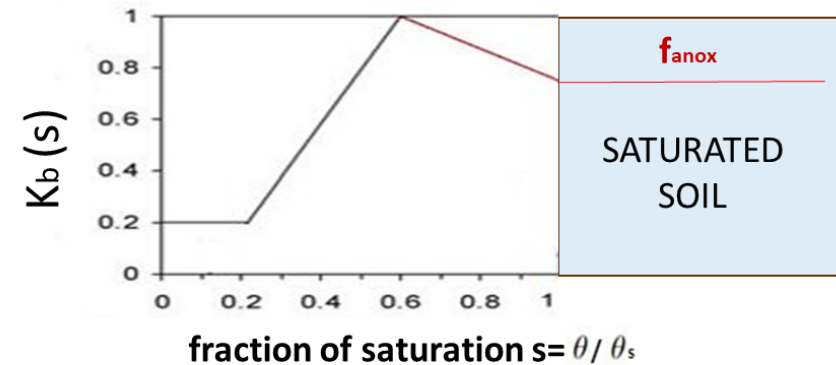


Model's assumption:

In large saturated zones of soil SOC decomposition is performed by different bacteria that respire CO₂ or CH₄



Zhang, Y., Li, C., Trettin, C. C., Li, H., & Sun, G. (2002). An integrated model of soil, hydrology, and vegetation for carbon dynamics in wetland ecosystems. *Global biogeochemical cycles*, 16(4), 9-1.



Modifying rate factor for soil moisture

$$K_b(s) = \begin{cases} 0.2, & \text{if } s(t) \leq s_{min} \\ 1 - 0.8 \frac{s_0 - s(t)}{s_0 - s_{min}}, & \text{if } s_{min} < s(t) \leq s_0 \\ 1 - (1 - f_{anox}) \frac{s(t) - s_0}{1 - s_0}, & \text{if } s_0 < s(t) \leq 1 \end{cases}$$

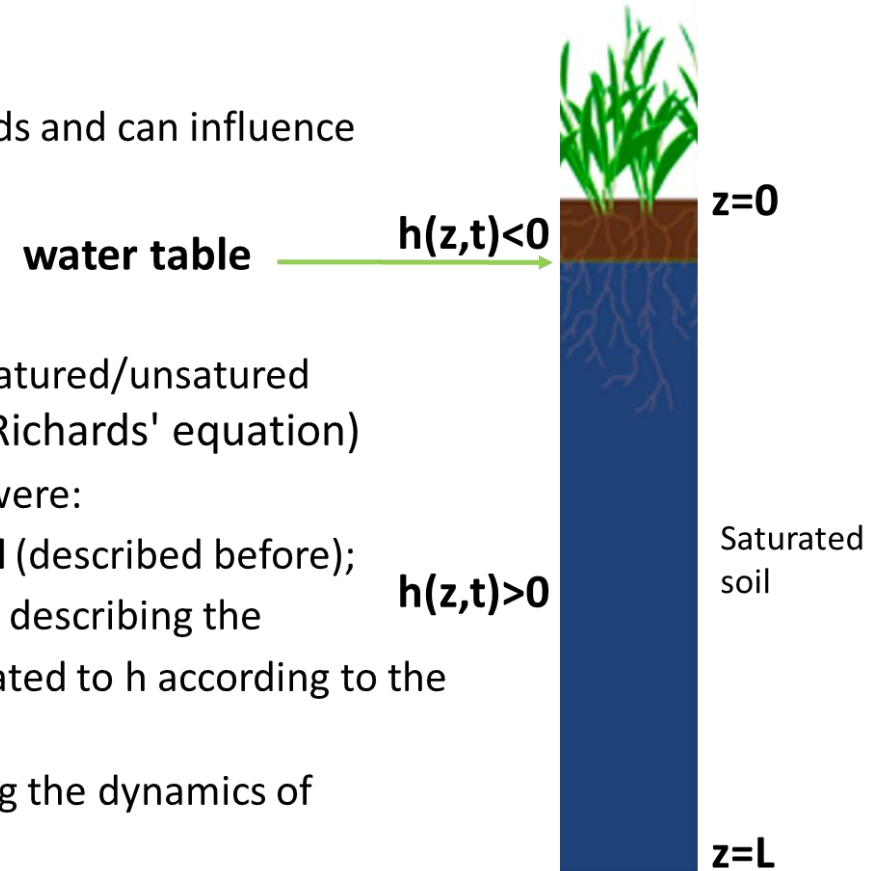
In-depth RothC for wetlands



The vertical water flow can be significant in wetlands and can influence the dynamics of in-depth SOC concentration

Model's structure

- Water head pressure $h=h(z,t)$ [L] dynamics in saturated/unsaturated soil described by **Richardson equation** (1D Richards' equation)
- **SOC diffusion-advection-reaction equations** were:
 - reaction given by **modified RothC for wetland** (described before);
 - advection term $v(z,t) = K(h) \left(\frac{\partial h}{\partial z} - 1 \right)$ [LT^{-1}] for describing the transport of SOC due to vertical water flow related to h according to the Darcy's law
- **temperature transport equation** for describing the dynamics of in-depth temperature distribution



In-depth RothC for wetlands

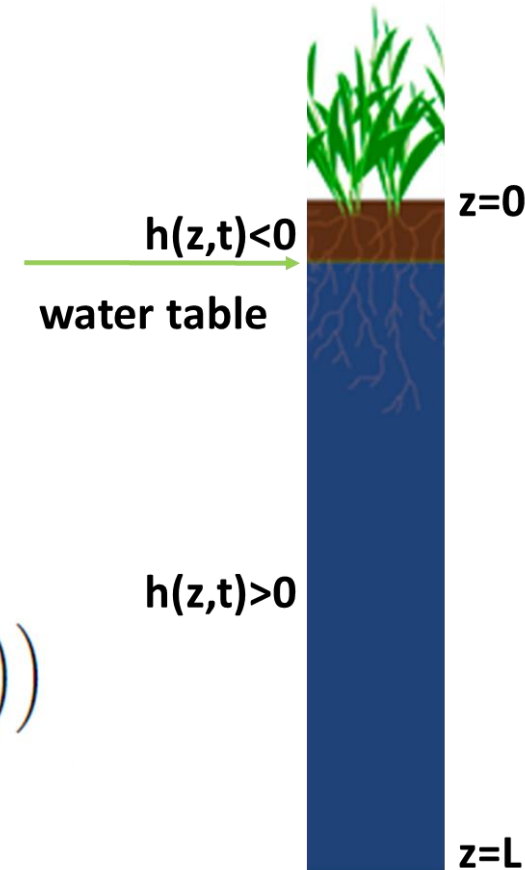


The vertical water flow can be significant in wetlands and can influence the dynamics of in-depth SOC concentration

Model's structure

- Water head pressure $\mathbf{h=h(z,t)}$ [L] dynamics
Richardson equation (1D Richards' equation)

$$\left(C(h, z) + \frac{\theta(h, z)}{\theta_s(z)} S_s(z) \right) \frac{\partial h(z, t)}{\partial t} = \frac{\partial}{\partial z} \left(k(h, z) \left(\frac{\partial h(z, t)}{\partial z} - 1 \right) \right)$$



In-depth RothC for wetlands



The vertical water flow can be significant in wetlands and can influence the dynamics of in-depth SOC concentration

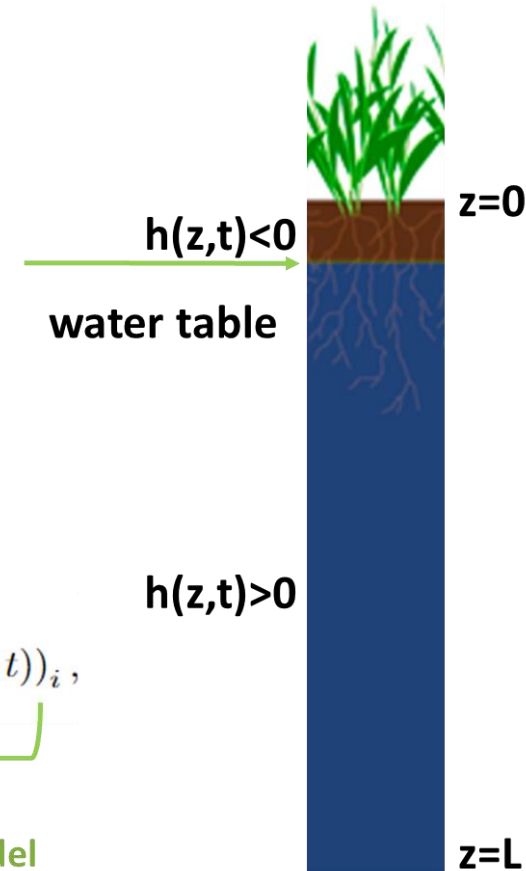
Model's structure

- SOC compounds diffusion-advection-reaction equations, the reaction part of which is in the form of RothC model for wetland;

$$\sigma \frac{\partial c_i(z, t)}{\partial t} = D_i \frac{\partial^2 c_i(z, t)}{\partial z^2} - v(z, t) \frac{\partial c_i(z, t)}{\partial z} + \underbrace{(\rho(z, t) A c(z, t) + b(z, t))}_i,$$

$$i = 0, \dots, 3$$

Original RothC model



In-depth RothC for wetlands



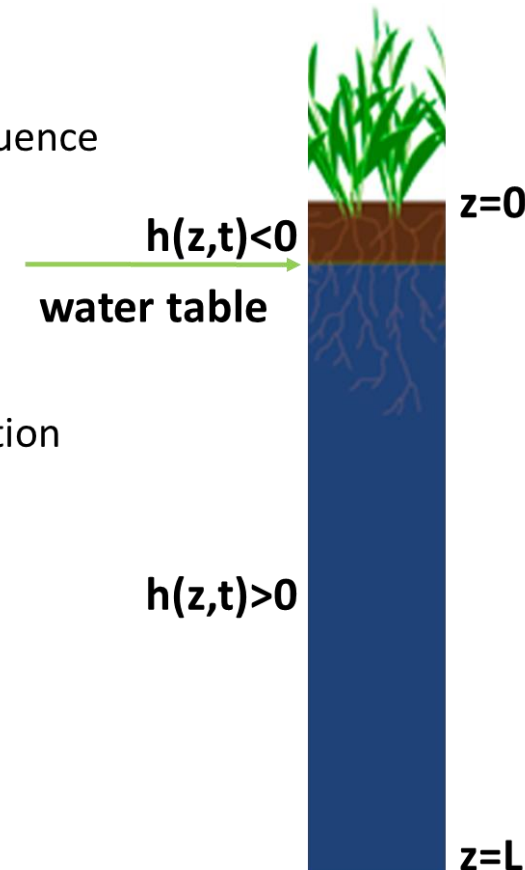
The vertical water flow can be significant in wetlands and can influence the dynamics of in-depth SOC concentration

Model's structure

- Temperature transport equation (to consider in depth distribution of temperature influence of microbial activity).

$$C_T \frac{\partial T}{\partial t} = \frac{\partial}{\partial z} \left(\lambda \frac{\partial T}{\partial z} \right) - v(z) \frac{\partial T}{\partial z}$$

$$T|_{z=0} = T_a, \quad \frac{\partial T}{\partial z} \Big|_{z=L} = 0$$



Boundary conditions



In-depth RothC for wetlands.

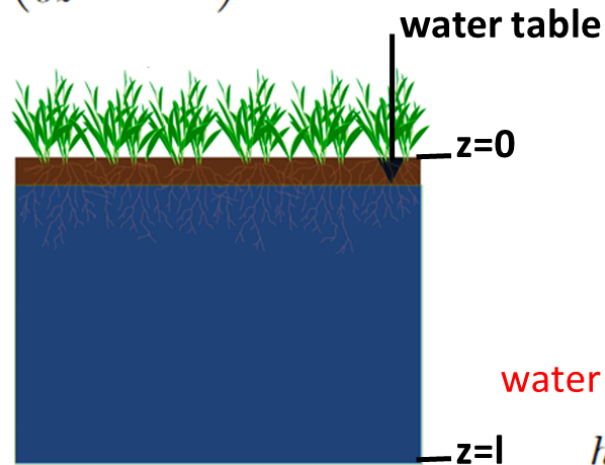
Boundary conditions that can take into account periodical flooding

$h(z,t)$: water head pressure

Non submerged soil:

prescribed flow (Neumann condition)

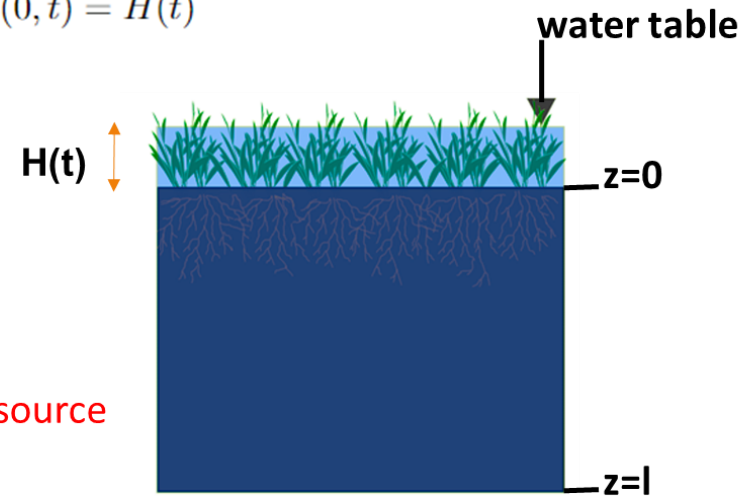
$$-K(h(0,t)) \left(\frac{\partial h}{\partial z}(0,t) - 1 \right) = P(t) - E(t)$$



Submerged soil:

prescribed value (Dirichlet condition)

$$h(0,t) = H(t)$$



$$l_e(t)$$

water level in an external source

$$h(l,t) = l - l_e(t)$$

Figures from: Asmuß, T., Bechtold, M., & Tiemeyer, B. (2019). On the potential of Sentinel-1 for high resolution monitoring of water table dynamics in grasslands on organic soils. *Remote Sensing*, 11(14), 1659.



The considered scenario was the growth of rice in the lands of Ebro Delta. Rice fields are flooded from the end of April to September-October. Flooding was modelled by linear change of water level $Le(t)$ from the value below bottom depth in December-January to the level of 10 cm above soil surface in May-September

Monthly averaged precipitation and air temperature data was taken from the website <https://en.climate-data.org/europe/spain/catalonia/amposta-56879/> for the city of Amposta, the nearest to the Ebro delta. Having only temperature data, evapotranspiration was calculated according to the Hargreaves-Saman formula

M. Belenguer-Manzanedo, C. Alcaraz, M. Martinez-Eixarch, A. Camacho, J. Morris, C. Ibanez, Modeling soil accretion and carbon accumulation in deltaic rice fields, Ecological Modelling 484 (2023) 110455

Parameters



Param.	Description	Value	Dimension	Ref.
θ_r	Residual water content	0.098		Obtained from clay, silt, and sand content using Rosetta v.1 model
θ_s	Saturated water content	0.422		-//-
a	van Genuchten model's parameter	0.005		-//-
n	-//-	1.436		-//-
K_s	Coefficient of filtration	$1.18 \cdot 10^{-2}$	m/day	-//-
β	Mualem model's parameter	0.678		-//-
S_s	Specific storage	10^{-5} – 10^{-3}	m^{-1}	Possible range of values for silt-rich soils according to the data in [5]
l	Bottom depth	1	m	[6]

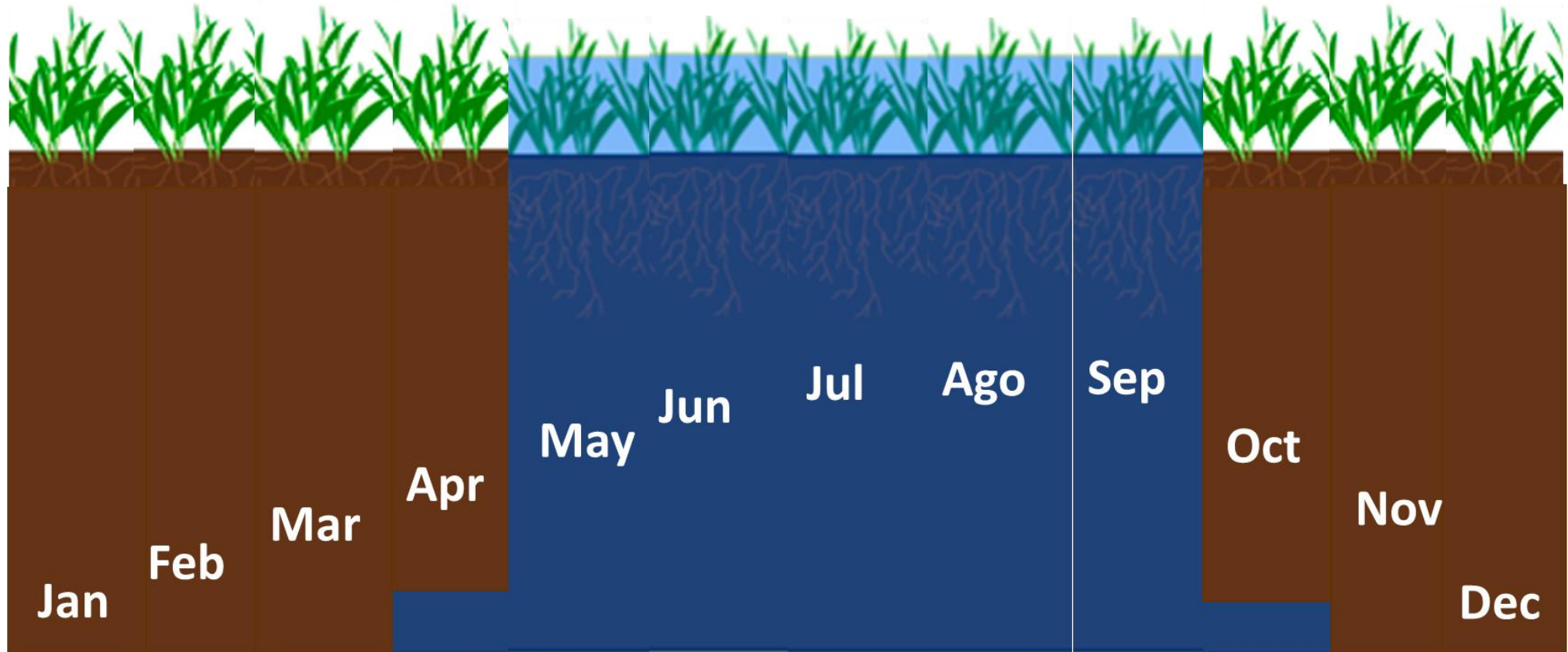
Table 1: The values of parameters for the Richards' - Richardson equation

Month	ET, mm/day	Precipitation, mm/day	L_e, m
Jan	0.95	1.47	4
Feb	1.21	1.12	2.5
Mar	1.81	1.38	1.5
Apr	2.59	1.81	0.75
May	3.37	1.90	-0.1
Jun	4.23	0.95	-0.1
Jul	4.41	0.77	-0.1
Aug	4.15	1.21	-0.1
Sep	3.28	2.33	-0.1
Oct	2.33	2.51	0.75
Nov	1.38	1.73	2.5
Dec	0.95	1.47	4

Table 2: The values of parameters



One-year simulation



References



Bohaienko, V.; Diele, F.; Marangi, C.; Tamborrino, C.; Aleksandrowicz, S.; Woźniak, E. A Novel Fractional-Order RothC Model. *Mathematics* 2023, 11, 1677. <https://doi.org/10.3390/math11071677>

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Diele, F., Luiso, I., Marangi, C. et al. Evaluating the impact of increasing temperatures on changes in Soil Organic Carbon stocks: sensitivity analysis and non-standard discrete approximation. *Comput Geosci* 26, 1345–1366 (2022), <https://doi.org/10.1007/s10596-022-10165-3>.

Diele, F., Marangi, C., Martiradonna, A. (2021). *Non-Standard Discrete RothC Models for Soil Carbon Dynamics*, *Axioms*, 10(2), 56.

<https://github.com/CnrlacBaGit/NSRothC>

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